

**CRA Appendices** 

EPA's Recertification Activities Webpage

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- 1 Fluid pressure above hydrostatic is a hydrologic characteristic of the Salado (and the Castile) that
- 2 plays a potentially important role in the repository behavior. It is difficult to accurately measure
- anatural pressures in these formations *accurately* because the boreholes or repository excavations
- 4 required to access the rocks decrease the stress in the region measured Stress released
- 5 instantaneously decreases fluid pressure in the pores of the rock so measured pressures must be
- 6 considered as a lower bound of the natural pressures Stress effects related to test location and
- 7 the difficulty of making long duration tests in lower permeability rocks result in higher pore
- 8 pressures observed to date in anhydrites The highest observed pore pressures in halite rich
- 9 units near Room Q are is on the order of 9 MPa whereas the highest pore pressures observed
- 10 in anhydrite are approximately 12 5 MPa (Beauheim et al. 1993-139 Beauheim and Roberts
- 11 2002, p 82) Far field pore pressures in halite rich and anhydrite beds in the Salado at the
- repository level are expected to be similar because the anhydrites are too thin and of too low
- permeabilities to have liquid pressures much different than those of the surrounding salt. For
- comparison the hydrostatic pressure for a column of brine at the depth of the repository is about
- 15 7 MPa and the lithostatic pressure calculated from density measurements in ERDA 9 is about
- 16 15 *MPa*
- 17 Fluid pressures in sedimentary basins that are much higher or much lower than hydrostatic are
- referred to as abnormal pressures by the petroleum industry where they have received
- 19 considerable attention In the case of the Delaware Basin evaporites the high pressures are
- almost certainly maintained because of the large compressibility and plastic nature of the halite
- 21 and to a lesser extent the anhydrite. The lithostatic pressure at a particular horizon must be
- supported by a combination of the stress felt by both the rock matrix and the pore fluid. In
- 23 highly deformable rocks the portion of the stress that must be borne by the fluid exceeds
- 24 hydrostatic pressure but cannot exceed lithostatic pressure
- 25 Brine content within the Salado is estimated at 1 to 2 percent by weight although the thin clay
- seams have been *inferred* observed by Deal et al. (1993 pp. 4.3) to contain up to 25 percent
- 27 brine by volume Where sufficient permeability exists this brine will move towards areas of
- lower hydraulic potential such as a borehole or mined section of the Salado
- Observation of the response of pore fluids in the Salado to changes in pressure boundary
- 30 conditions at walls in the repository in boreholes without packers in packer sealed boreholes or
- 31 in laboratory experiments is complicated by low permeability and low porosity. Qualitative data
- 32 on brine flow to underground workings and exploratory boreholes have been were collected
- routinely between since 1985 and 1993 under the Brine Sampling and Evaluation Program
- 34 (BSEP) and have been documented in a series of reports (Deal and Case 1987 Deal et al. 1987
- 35 1989 1991a 1991b and 1993, and 1995) These and other investigations are discussed in
- 36 Appendix SUM (Section 3 3 1 3) A discussion of alternative conceptual models for Salado fluid
- 37 flow is given in Appendix **PA**, Attachment MASS Section MASS 7 Additional data on brine
- 38 inflow are available from the Large Scale Brine Inflow Test (Room Q) Flow has been observed
- 39 to move to walls in the repository to boreholes without packers and to packer sealed boreholes
- 40 These qualitative and relatively short term observations suggest that brine flow in the fractured
- 41 DRZ is a complex process In some locations evidence for flow is no longer observed where it
- 42 once was in others flow has begun where it once was not observed. In many cases
- observations and experiments must last for months or years to obtain useful results

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