

DESERT PAVEMENT

Desert pavement is a thin, continuous layer of closely spaced, stones or pebbles that cover the surface of the soil. It is formed by the accumulation of material that has been transported from the mountains and deposited in the valleys. The pavement is composed of a variety of materials, including sand, silt, and clay, and is typically 1 to 2 inches thick. It is a common feature of the desert floor and is often found in association with other desert features, such as dunes and washes.

CAVE DEPOSITS

Cave deposits are formed by the accumulation of material that has been transported from the mountains and deposited in the valleys. The deposits are composed of a variety of materials, including sand, silt, and clay, and are typically 1 to 2 inches thick. They are often found in association with other desert features, such as dunes and washes.

DESERT TERRACES

Desert terraces are formed by the accumulation of material that has been transported from the mountains and deposited in the valleys. The terraces are composed of a variety of materials, including sand, silt, and clay, and are typically 1 to 2 inches thick. They are often found in association with other desert features, such as dunes and washes.

DESERT VARNISH

Desert varnish is a thin, continuous layer of closely spaced, stones or pebbles that cover the surface of the soil. It is formed by the accumulation of material that has been transported from the mountains and deposited in the valleys. The varnish is composed of a variety of materials, including sand, silt, and clay, and is typically 1 to 2 inches thick. It is a common feature of the desert floor and is often found in association with other desert features, such as dunes and washes.

TRANSITIONAL DEPOSITS

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TRANSPORTED DEPOSITS

Transported deposits are formed by the accumulation of material that has been transported from the mountains and deposited in the valleys. The deposits are composed of a variety of materials, including sand, silt, and clay, and are typically 1 to 2 inches thick. They are often found in association with other desert features, such as dunes and washes.

ALLUVIUM IN FLOODPLAINS AND STREAM CHANNELS

Alluvium in floodplains and stream channels is formed by the accumulation of material that has been transported from the mountains and deposited in the valleys. The alluvium is composed of a variety of materials, including sand, silt, and clay, and is typically 1 to 2 inches thick. It is a common feature of the desert floor and is often found in association with other desert features, such as dunes and washes.

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GRAVEL TERRACES — Well-rounded stream gravels with cobbles 6 inches or more in diameter, some terraces 250 ft higher than the streams. Especially well developed along the San Juan River, less so along the Pecos, Gila, and Canadian Rivers. Most recent deposits by Pleistocene melt waters from mountains. Abundant caliche deposits, especially on the higher terraces, which may be Kansan; lowest are Wisconsinan.

ALLUVIAL FAN DEPOSITS

In alluvial fans, unlike floodplain alluvium, beds tend to be thick, massive, and highly lenticular rather than well stratified. This is characteristic of all the facies, whether boulder, gravel, sand, or silt. Beds lenticular and elongated down the slope of the fans; slopes 2 to 20 percent. Deposition mostly by flash floods, with poor sorting and mixed textures. Coarse-textured lenses commonly form ridges extending down the fan onto generally finer grained sediment. Boundaries between the textural facies of the deposits roughly parallel the fan contour, but detailed boundaries are irregularly lobate; those shown are approximations. Fan textures and slopes depend partly on composition of the parent rocks and partly on height and steepness of the bordering hill or mountain. Fans extensive in the Basin and Range part of the state where they comprise about half the total area; in other parts of the state, fans are small. On the larger fans, arroyos become shallower towards the toe, many head at low mounds that probably mark old mudflows. Ground subject to sheet flooding.

GRAVEL FACIES — Boundary towards apex of fan, grading downward to cobble and fine gravel with increasing proportion of sand and finer grained material. Commonly dissected to form 2 to 3 levels of gravel benches up to 50 ft above present washes. A few streams (e.g., Mulligan Wash, Alamosa River, Cuchillo Negro Creek, and Rio Canon Arroyo) are incised 100 ft below fan surfaces. On short, steep fans, depths of valleys generally decrease downslope. On the broad Palomas surface, west of the Rio Grande above Hatch, valleys maintain their depth. Except near the apex, extensive surfaces have smooth desert pavement. On short, steep fans, gravel shows minimal weathering and are weakly cemented with caliche, age probably Wisconsinan and Holocene. On broad, more gently sloping fans, gravels are more weathered and commonly cemented by caliche, age probably pre Wisconsinan. In south half of the state, gravel facies is characterized by creosote bush cover. Thin alluvial gravel covering pediments is denoted by Ig over subscript that identifies parent formation.

SAND FACIES — Sandy alluvium with subordinate amounts of fine gravel, silt, and clay. Forms at least four kinds of ground: 1) On short, steep fans sloping from the mountains of granitic or gneissic rock (e.g., parts of the Florida Mountains), this facies may form a smooth sandy layer a few feet thick covering gravel below; slopes 5 to 20 percent; washes 1 to 10 ft deep may expose underlying gravel. 2) On other short fans, sand facies may form arcuate belt at toe of fan with slopes averaging 10 percent; commonly reworked into coppice dunes 3 to 7 ft high (sm). 3) Other belts of smooth sandy ground commonly slope 5 percent or less and consist of sand mounds approximately 1 ft high over caliche (ls). 4) Gypsiferous sand (ls_g), especially in the Jornada del Muerto, Tularosa Valley and east side of the Rio Valley. Sand facies absent on the broad Las Palomas surface. Thin fan sand covering pediments is denoted by Is over subscript that identifies underlying formation. Boundary with residual sand, fan gravel, and fan silt is approximate.

SILT FACIES — In Basin and Range parts of the state, toes of fans may be silty and clayey rather than sandy; surface smooth, with slopes less than 5 percent. Slow infiltration of rain and low slopes result in sluggish runoff. Forms a belt below the sand facies and grades downward to plays silt (psl) with slopes less than 2 percent. Abundant swelling clays and exchangeable sodium. Surface layers predominantly Holocene; subject to sheet flooding, preglacial with als. East and west of Sangre de Cristo Mountains, also forms fans of sandy or silty loam with little gravel in upper 3 to 4 ft, but abundant gravel below the loam. Caliche soft. Includes loess on isolated hillslopes. Boundary with residual loam (rl), plays silt (psl), and fan sand (Is) approximate.

FOLIAN DEPOSITS

Eolian deposits are laid down by wind, mostly as sheets of sand or silt (loess). Rarely, after prolonged drought on shale desert in the San Juan Basin, shale flakes may accumulate in rippled sheets or even small dunes, but with the next rain, these become mud. Sand dune shapes depend on topography, relative strength of the winds, supply of sand, and vegetation. Some dunes are concave towards the windward (parabolic), others are concave towards the leeward (barhans), and others are longitudinal or transverse (line dune clusters (e.g., Great White Sands) have all four kinds. Dunes may climb a windward slope or fall on a leeward slope. Most of New Mexico's eolian sand sheets have a basal layer of weathered, partly cemented, reddish stabilized sand, some sand surfaces on such layers are smooth, in the Basin and Range and Great Plains parts of the state, these surfaces are generally underlain by caliche, in the San Juan Basin, sand sheets commonly overlie residual, fan deposits, or bedrock. Where sand is thick, as on sand facies of fans in the Basin and Range and at climbing dunes east of the Pecos River (Mesquero Sands) the sand is in mounds (coppice dunes) with profuse growth of vegetation, mesquite, and saltbrush in the Basin and Range, sand sage, thorny oak, small soapweed vine, and occasional mesquite on the Mesquero Sands. Sand sheets are predominantly late Pleistocene, mounds and dunes are largely Holocene.

SAND UNDERLAIN BY BASALT — Extensive on basaltic plains south and east of Zuni Mountains and on West Potrillo Mountains. At Kiltwater Hole and Hunt's Hole, the sand is of volcanic origin.

SAND UNDERLAIN BY CALICHE ON SANTA FE GROUP — Mostly on La Mesa and south part of the Jornada del Muerto.

THIN SAND ON CALICHE ON OGALLALA FORMATION — Thickness about 1 ft. Chips of caliche comprise 30 percent of the sand. Generally too shallow for farming, but good shallow source for aggregates.

MODERATELY THICK SAND ON CALICHE ON OGALLALA FORMATION — Sand 1 to 3 ft thick. Surface layers noncalcareous over reddish loam. Local sand mounds. Ground favorable for farming. Boundaries approximate.

THICK SAND ON CALICHE ON OGALLALA FORMATION — Sand 3 to 5 ft thick. Local mounds. Brownish-red, fine sandy loam over reddish brown, sandy clay loam, noncalcareous. In depths of 3 ft, calcareous subsoil contains filaments of lime carbonate. Where farmed, ground is subject to wind erosion. Boundaries approximate.

LOOSE SAND IN MOUNDS — Coppice dunes, commonly 2 to 7 ft high and 25 to 50 ft in diameter; generally elongated north of east but a local exception lies east of Columbus where elongation is south of east. Age is Holocene. Boundaries fairly accurate.

SAND SHEETS — Surfaces smooth except for ripples 2 to 3 inches high and scattered sand mounds 3 to 12 inches high, especially around small shrubs. Thickness of loose sand generally no more than about 12 to 24 inches, but commonly overlies stabilized sand. Underlying material where known, identified by subscript.

LONGITUDINAL DUNES — Sand commonly 6 ft thick, locally 10 ft. Forms distinct ridges generally oriented north of east. Locations diagrammatic and width exaggerated.

OTHER DUNES — ds₁, quartzose sand; ds₂, gypsiferous sand.

LOAM ON OLD BASALTIC LAVA — Probably pre Wisconsinan loess.

FOLIAN SILT

EXPLANATION OF SURFICIAL GEOLOGY

by Charles B. Hunt 1977